



end users report: Potentially damaging offshore storm conditions along the Sefton Coast.



By Jenny Brown



- Introduce the study area
- Modelling methods
- •Show a past storm event used to validate the model setup
- •Present the results from an 11year hindcast
- Historical events
- Data use / Conclusions











Crosby

Formby storm 31<sup>st</sup> March – 1<sup>st</sup> April 2010

## Natural defence – risk of erosion

Formby extreme high tide Feb 2010

#### **Provide Important Habitats**



#### **Economic Value**



## Coastal Management requires the best knowledge of storm conditions



Develop state-of-the-art wave-tide-surge Liverpool Bay modelling system

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#### Wave Modelling: WAveModel

State of the art 3<sup>rd</sup> generation spectral wave model – extended for shallow water



## Tide – Surge Modelling: ProudmanOceanographicLaboratoryCoastalOceanModellingSystem

3D hydrodynamic model – tidal, riverine and meteorological forcing



## Irish Sea: Wave-tide-surge interaction

Exchange information between surge model and wave model



Coupling involves:

- (I) time varying depth and velocity fields, which refract the waves
- (II) wave-current bottom boundary layer
- (III) wave dependent surface roughness to generate the surge

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#### 18<sup>th</sup> January 2007

#### Depression travelled east to the north of Ireland and across Scotland.



#### 11-12<sup>th</sup> November 1977

Low pressure system moved from the west, easterly over the north of Scotland.



#### **Location Map**





#### **Model Surge Results 1977**



#### No Wave data available

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Observation + model hindcast at tide gauges either side of Sefton Coast

#### Generated: SSW – WSW winds (longest fetches)

The external surge from the SW approaches/Celtic Sea dominates the surge generation.

Reach: 2.6m

#### Flood Risk in the study area



Large 10m tidal range Liverpool Bay causes tide-surge interaction, preventing the peak surge at HW.



Greatest risk SW wind veers W during spring tides.

-250 Extreme wave generation occurs on top of extreme high tide levels, which are increased by extreme surge conditions.



# Changes in the Dune system

Dune toe position surveyed on the 24<sup>th</sup> September 2002 and retreat following this survey until the 26<sup>th</sup> November 2002

Large changes around the point – expose (new) areas to flood risk during erosion period and during following high tide before dune recovery.

Retreat due to:

- (i) storm event 25<sup>th</sup> -27<sup>th</sup> October 2002 SW-W winds, 2.3m surge, 4m waves & consecutive tides >8.7m CD.
- (ii) Sequence of extreme high tidal levels during the autumn period (spring tide & any wind event SW-W).



![](_page_27_Figure_0.jpeg)

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![](_page_28_Picture_7.jpeg)

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![](_page_28_Picture_9.jpeg)

Historical events Often associated with veering SW-W wind

Breaching of sea wall at Southport, May 1960 (Image courtesy of SMBC).

Past damaging storm events to the Sefton coast reported in: 1961, 1965, 1967, 1968, 1975, 1976, 1977, 1983, 1990, 1997, 2002, 2004 & 2008

> Flooding at Marine Drive, Southport, November 1993 (Image courtesy of SMBC)

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![](_page_30_Picture_7.jpeg)

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![](_page_30_Picture_9.jpeg)

#### Conclusions

•POLCOMS-WAM is a valid wave-tide-surge model to apply to the Irish Sea and Liverpool Bay.

Can now be used for future flood risk projection.

•Extreme surge levels (2.6m) along the Sefton coast occur due to SW winds.

 Extreme wave heights (5.6m offshore / 2.5m nearshore) in Liverpool Bay occur in response to NW – W winds.

•Storm tracks generating SW winds veering W lead to most severe conditions. E.g. the historical Nov 1977 event.

![](_page_31_Picture_6.jpeg)

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